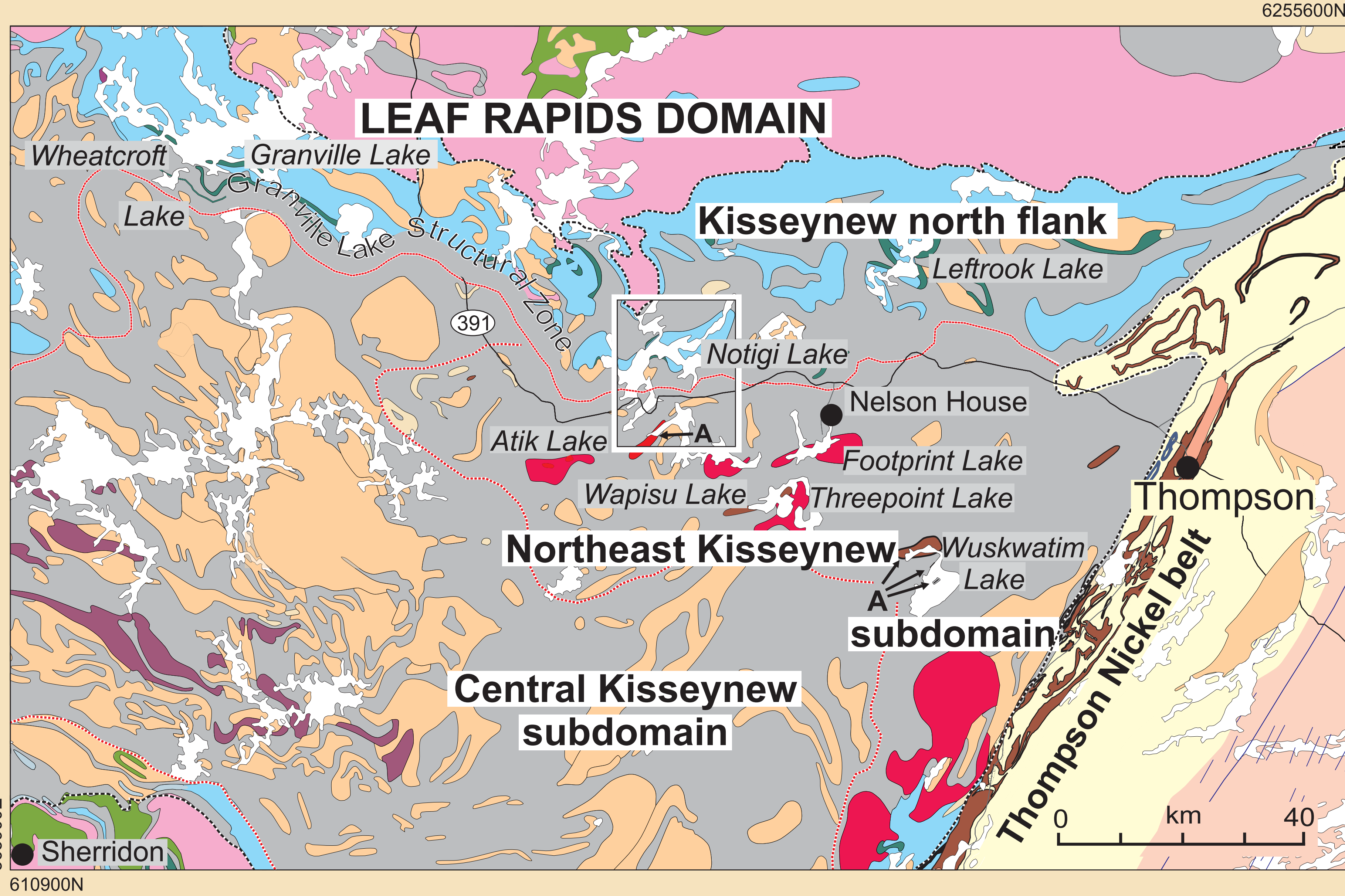


# Stratigraphy and Structure of the Notigi–Wapisu Lakes Area

**Manitoba** by H.V. Zwanzig, L.A. Murphy and N. Rayner

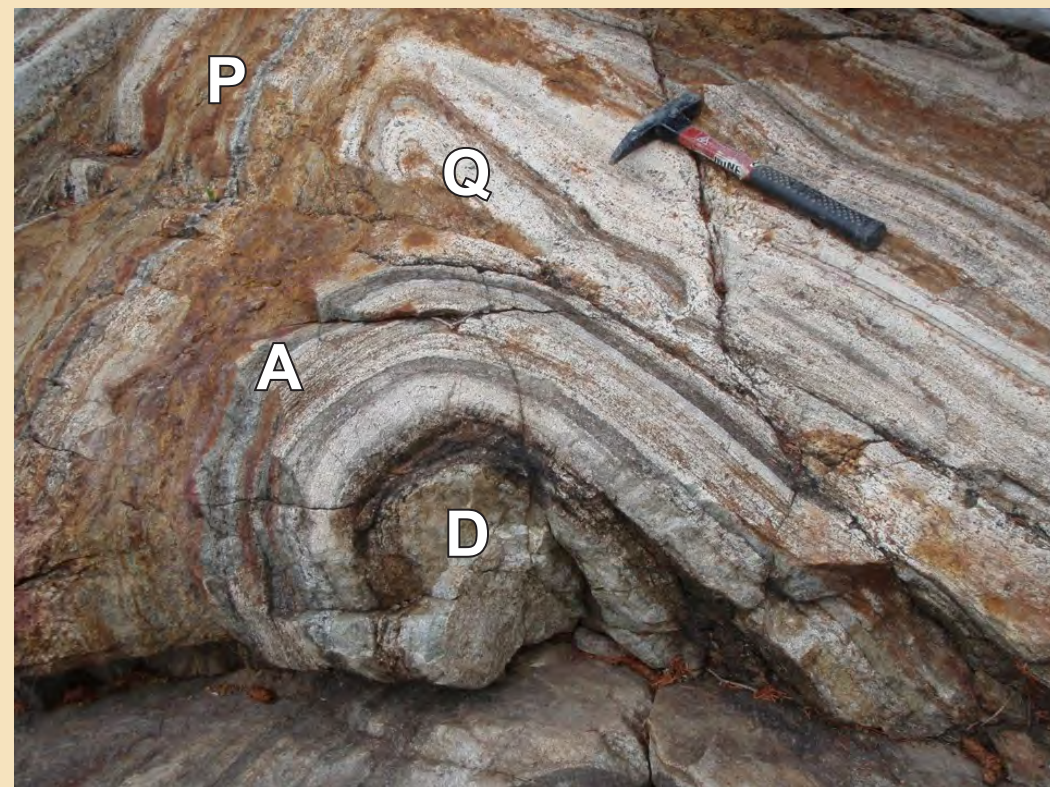
**Introduction** The Notigi-Wapisu lakes area (white outline below, and map on right) provides a transect from the south side of the Leaf Rapids Domain across the Kisseynew north flank into the Northeast Kisseynew subdomain. These areas represent, respectively, a 1.90–1.84 Ma volcano-plutonic arc terrane with VMS potential, a 1.85–1.83 Ga basin margin with sedimentary Cu potential but also containing a suture zone (Granville Lake structural zone) with Au potential, and 1.85–1.83 Ga basinal sedi-mentary rocks interleaved with pericontinental Archean basement with cover rocks that have potential for Thompson-type Ni deposits.



**Tectonostratigraphy** can be restored to an early stage of thrusting and folding using the regional relationships on the Kisseynew north flank. Sedimentary facies change from distal turbidite (Burntwood Group) to various proximal arkosic compositions (Sickle Group). A thin intervening structural package (Granville Lake assemblage) includes older metavolcanic and metasedimentary rocks thrust over the Burntwood Group with the Sickle Group lying unconformably on top (*see* column on right corresponding to section A–A' on geologic map). The older rocks on Wapisu Lake (Wuskwatim Lake sequence) include a 1.5 m layer and a 3.5 m layer of dark grey intermediate paragneiss separated by 1.6 m of semipelitic gneiss and followed by a 2 m section of sulphidic semipelite and silicate facies iron formation (hornblende-garnet bearing) and 2 m of amphibolite. Local protoquartzite establishes an Archean provenance for the sequence (*see* **Geochronology**). Contacts with the Burntwood Group are interpreted as faults and with pre-Burntwood quartz monzonite as intrusive but obscured by younger leucogranite (*see* column at far right).



**Wuskwatim Lake sequence** sulphidic semipelite. Note that similar looking rock in the Granville Lake assemblage (photo to right) is juvenile (Paleoproterozoic arc-derived).



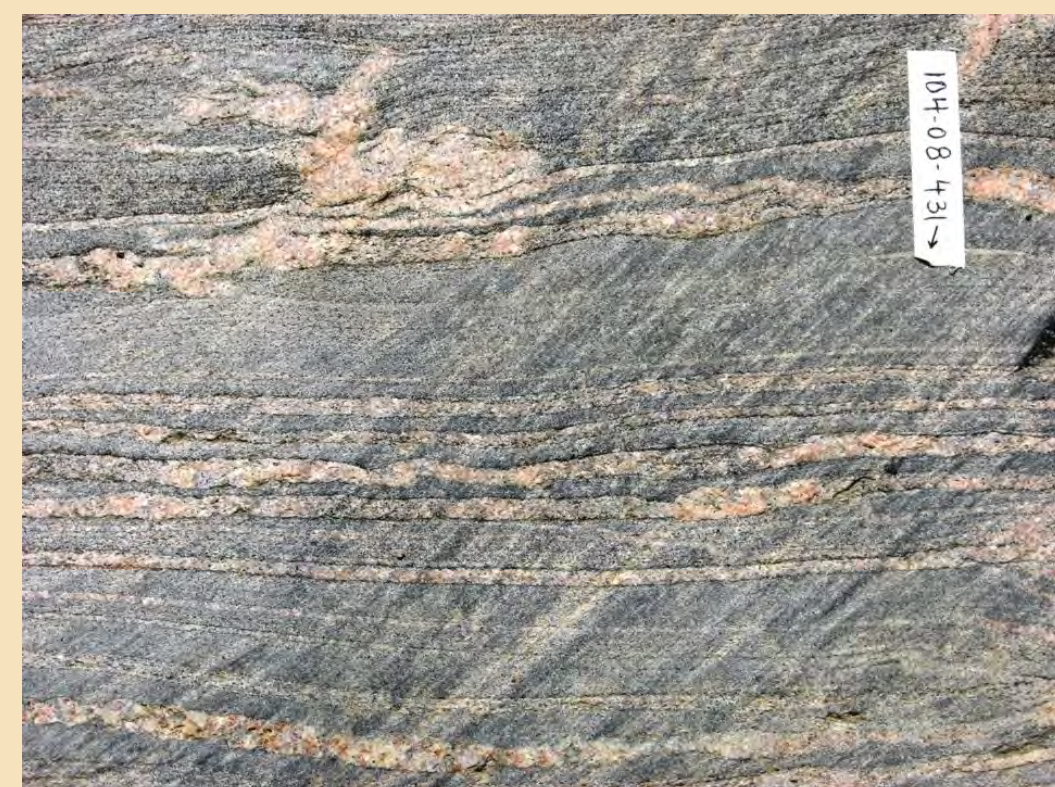
**Granville Lake assemblage** with interlayered sulphidic pelite (P), quartzofeldspathic to calcisilicate gneiss (Q), amphibolite (A) and diopside rich rock (D), interpreted to have been derived from mafic volcanic and sedimentary rocks. The fold is F<sub>3</sub>, showing a flexural style with a small thrust.



**Burntwood Group** garnetiferous metagreywacke-mudstone-derived metatextite.



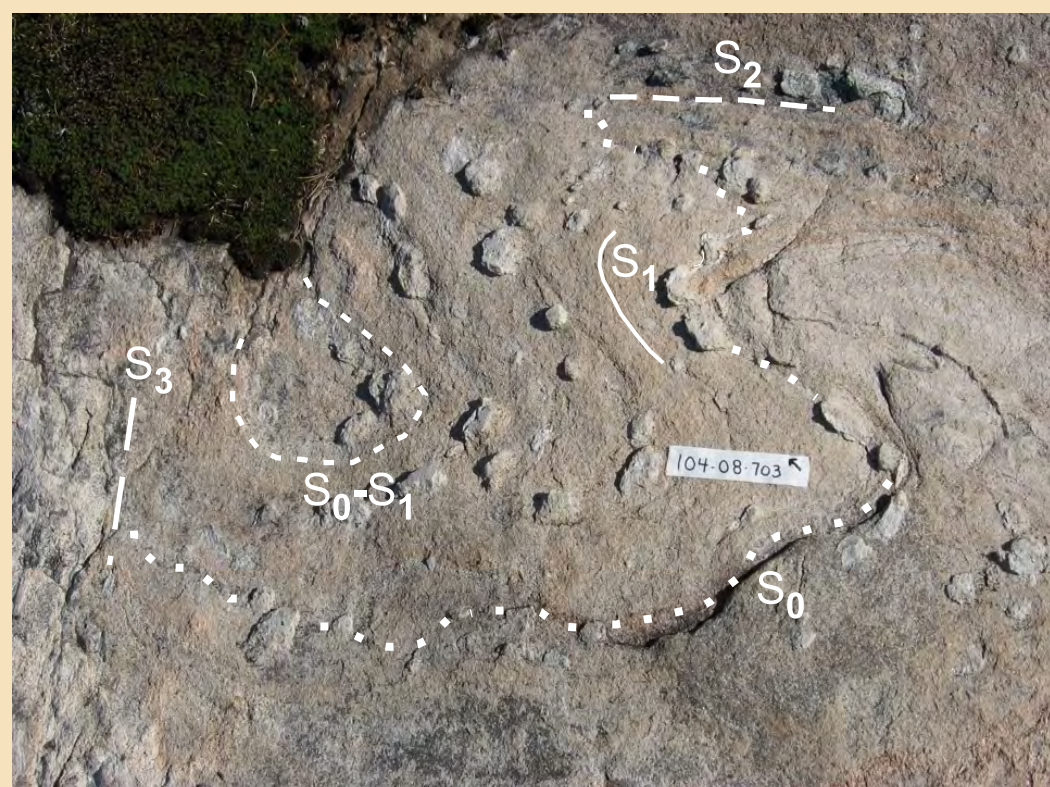
**Burntwood Group** metatubidite grading from greywacke (pale) to mudstone (darker) with tops shown as arrows.



**Sickle Group** magnetite-bearing biotite gneiss probably derived from nonmarine lithic arenite (redbeds).



**Sickle Group** hornblende-biotite gneiss (hornblende indicated by arrows) derived from arenite with probable calcite cement.



**Sickle Group** sillimanite-biotite gneiss (meta-arkose) with quartz-sillimanite knots; tape is along the axial plane of an F<sub>3</sub> fold and fabric elements are interpreted as labelled; note overprinting of F<sub>2</sub> by S<sub>3</sub>



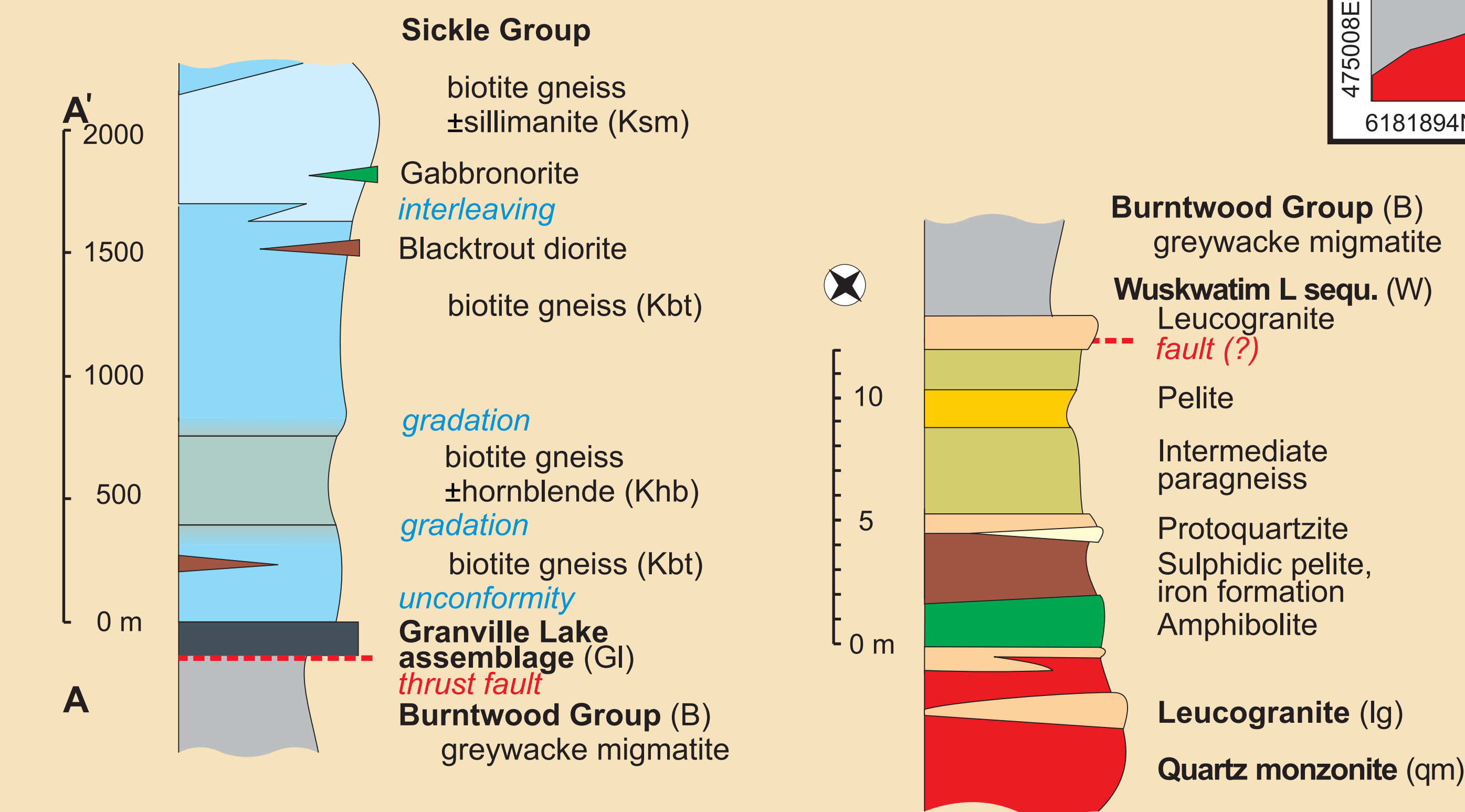
**Isoclinal fold** in greywacke (lighter coloured outlined eye shape) interpreted as F<sub>1</sub> affected by later folds and pinch-and-swell structure; synkinematic granitoid veins show F<sub>2</sub> folding and injection into the axial plane.



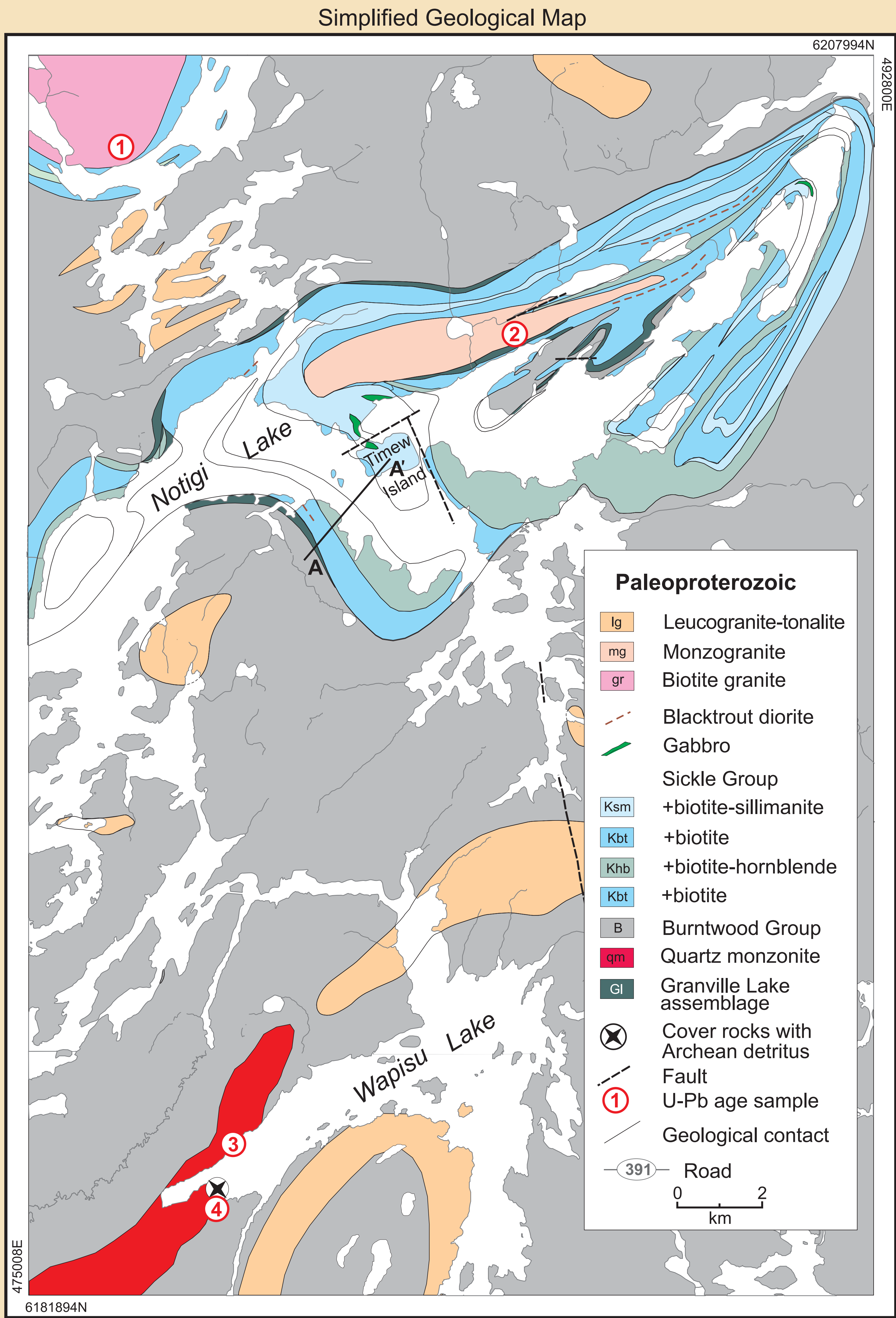
**Recumbent F<sub>2</sub> folds** of bedding, early foliation and granitoid veins as seen on a steep rock face.



**F<sub>2</sub>-F<sub>3</sub> interference structures** near the hinge zone of a major F<sub>3</sub> anticline. Note similar map-scale pattern at Notigi Lake.



**Structure** at Notigi Lake was interpreted using map patterns and changes in plunge produced during polyphase folding (map on right). The geometry was analyzed in 3-D using subareas with similar plunge to construct down-plunge sections with views that are nearly orthogonal to three different sets of folds (figures on far right). The Notigi (main) dome was formed by the interference of a late upright antiform (F<sub>3</sub>) with complex midcrustal recumbent structures (F<sub>2</sub>) that refolded major early folds (F<sub>1</sub>) and a thrust. Applying this history to the economically important rocks in the Northeast Kisseynew subdomain indicates that this area, although not widely underlain by thick Archean basement at depth, hosts narrow structural inliers many kilometres long similar to rocks in the Thompson Nickel Belt that host the nickel deposits.



**Uranium-lead geochronology** was carried out using the LA-MC-ICMS at the University of Alberta (**1**) and the SHRIMP at the Geological Survey of Canada in Ottawa (**2–4**). Previously obtained ages are on (**5**). The Rat pluton (**1**) is a foliated to gneissic, medium-grained biotite granite. It forms an asymmetric dome or sheath fold surrounded by the Sickle Group at the south margin of the Leaf Rapids Domain. Its crystallization age of 1838 ±4 Ma is similar to granodiorite at the same margin 50 km farther east. The age is considered as syndepositional with the Sickle Group, the lower part of which is apparently intruded. With plutons of known ages of 1.86 and 1.89 Ga also at the south margin, the Rat granite indicates that arc magmatism there spanned the same time period as throughout the internal (volcanoplutonic) zone of Trans-Hudson orogen. The Notigi granite (**2**) is a uniform foliated monzogranite near the core of the Notigi dome where it is interpreted to intrude the Sickle Group. Its zircons have distinct 1862 ±6 Ma cores and 1806 ±9 Ma rims (**2a, 2b**) but its field relationship indicates an age of emplacement younger than the 1850 Ma maximum age of deposition of the Sickle Group. Consequently, the cores are interpreted to be inherited or else the pluton was remobilized during high-grade metamorphism. The Wapisu quartz monzonite to quartz syenite (**3**) forms an elongate pluton assigned to the Footprint Lake plutonic suite of Whalen et al. (2008)<sup>1</sup>. Because of its better exposure but close relationship to the economically important Wuskwatim Lake sequence (*see* below), an early age would make this pluton another important probe for Archean rocks and their cover similar to bodies toward the east and west. This interpretation is confirmed by the crystallization age of 1872 ±9 Ma.

The Wapisu quartzite (**4**) is interpreted as part of the Wuskwatim Lake sequence which is related to the Ospwagan Group in the Thompson Nickel Belt where those rocks cover Archean basement and host all the nickel deposits. The 2.7 Ga peak is a common age in the Superior craton and the Mesoproterozoic peaks are interpreted to be derived from Assean-type crust at the north margin of the craton. All the Kisseynew sediments, in contrast, have detrital zircon ages that peak at 1855 Ma (**5**). They are clearly distinguished by their predominantly Proterozoic arc provenance from similar looking rocks in the Wuskwatim Lake sequence (*compare* plots on right).<sup>1</sup>MGS Report of Activities

