#### Project overview

Rocks of the Snow Lake - Squall Lake - Herblet Lake area include a panel of Snow Lake arc assemblage volcanic rocks imbricated between the McLeod Road thrust and the Birch Lake Fault. This panel is host to the most significant gold deposits, which together are responsible for production in excess of 1.4 million ounces of gold.

The rocks in the hangingwall of the McLeod Road thrust fault and farther north have only been mapped at 1:50 000 scale. In order to upgrade the geological map of the Snow Lake-Squall Lake-Herblet Lake area to a level comparable to that for the main panel of Snow Lake arc assemblage rocks to the south, the Manitoba Geological Survey is conducting new geological mapping. A particular emphasis of the project is the investigation of the geometry and stratigraphy of the New Britannia mine horizon in order to better understand the controls on gold

A new 1:20 000 geological map (PMAP2010-3) has been produced, integrating data from this latest field season with results from the previous years' mapping.

## Geology

The New Britannia mine horizon is part of the McLeod Road - Birch Lake thrust panel (MBP) of Snow Lake arc assemblage volcanic rocks. The MBP comprises a 2-3 km thick sequence of bimodal arc-related volcanic and volcaniclastic rocks that form a thin arcuate package extending more than 10 km east and west of Snow Lake. The rocks were metamorphosed from lower to mid- amphibolite facies. Synvolcanic gabbro dikes and sills are common. The MBP is flanked to the north and south by younger metasedimentary rocks. The MBP is separated from the younger underlying Burntwood metaturbidites by the McLeod Road thrust. The Birch Lake fault, also interpreted as a thrust fault, marks the boundary between the MBP and the overlying younger Missi meta-arenite to the north. The western portion of the MBP comprises two panels of pillowed and massive mafic volcanic flows with ocean-floor signature.

The northeastern portion of the map area, north of Herblet Lake between Angus Bay and Northeast Bay, is characterized by the presence of a large domal structure, the Herblet gneiss dome, exposing granodioritic to tonalitic gneiss in its core. This gneiss dome is part of a series of domes that extends more than 80 km west from the Pulver gneiss dome, just east of Herblet Lake, to the Sherridon–Hutchinson Lake complex.

# Updated 1:20 000 Scale Geological Map for the Squall Lake–Snow Lake–Herblet Lake Area

hornblende alteration along pillow

selvages and between pillows.

matrix of intermediate composition.

with thin selvages and rare interstitial

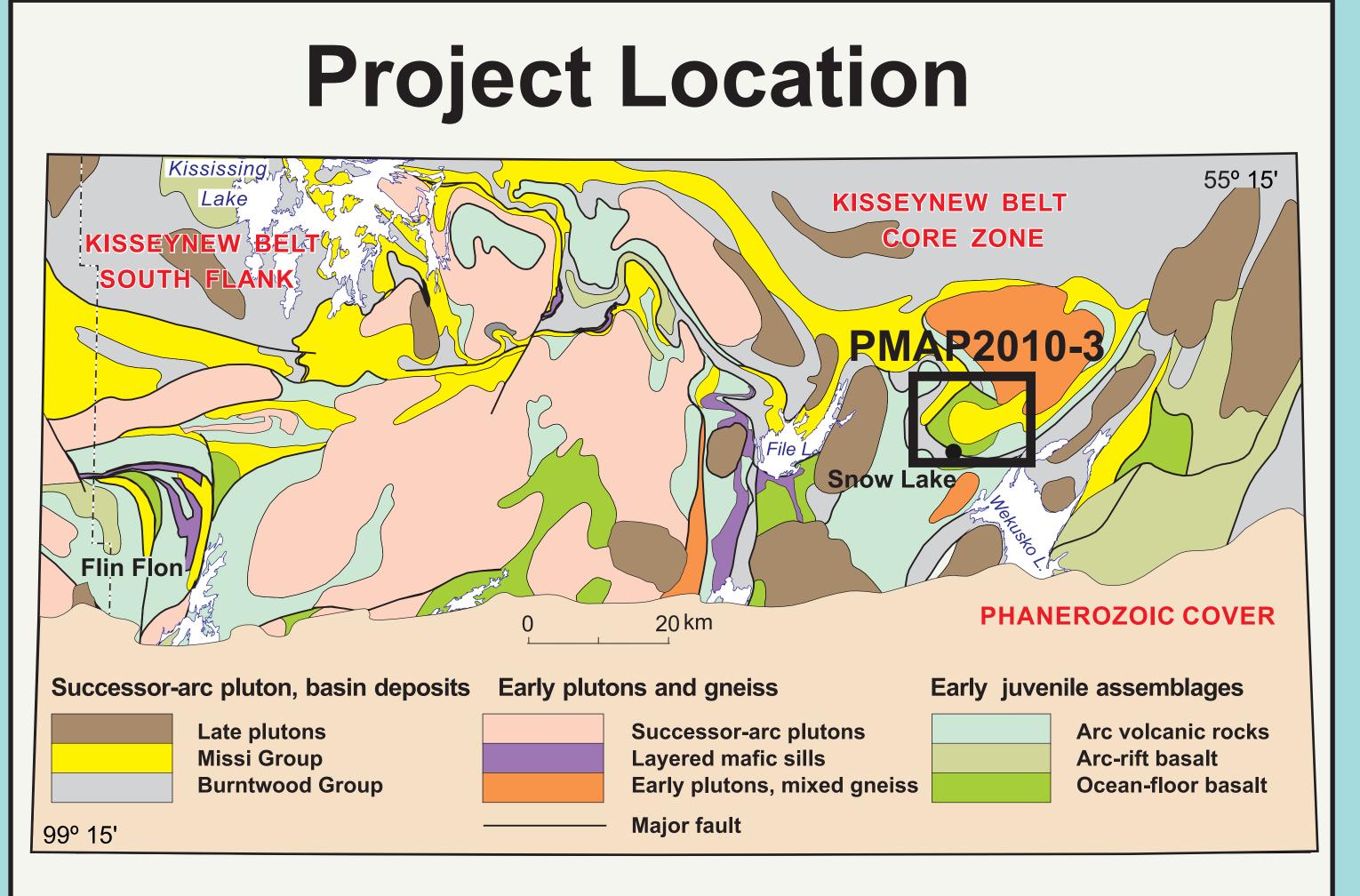
material.

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and crystal tuff.

(staurolite-rich) and psammitic beds.



Map showing the main geological subdivisions of the Flin Flon - Snow Lake greenstone belt. The project area is highlighted by the red box. The Snow Lake area is dominated by bimodal arc-related volcanic and volcaniclastic rocks. Younger Burntwood and Missi sediments are tectonically interleaved with the volcanics.

## Structural Geology

The Snow Lake – Squall Lake –Herblet Lake area was affected by four deformations (Kraus, 1998; Kraus and Williams, 1999) but only three generations of fabric elements representing deformations D<sub>1</sub>–D<sub>3</sub> have been identified.

The oldest deformation (D<sub>1</sub>) for which fabric elements are preserved in rocks of the map area produced upright to moderately inclined isoclinal folds and an associated axial-planar slaty cleavage. The presence of F₁ folds is widespread but the penetrative development and preservation of S₁ is rare. The isoclinal geometry of F<sub>1</sub> folds resulted in the development of a layer-parallel foliation, which was either not penetratively developed or was developed but transposed during subsequent deformations. Layer-parallel foliation of the S<sub>1</sub> generation is also observed in the hinge zone of tight to open, mesoscopic F<sub>2</sub> folds where S<sub>2</sub> overprinting did not develop as strongly. Throughout the map area, there is no evidence for the presence of an S<sub>1</sub> foliation within the Missi Group clastic sedimentary rocks. Within the New Britannia mine horizon isoclinal F. fold intercalation of volcanic rocks was documented. The plunge of F<sub>1</sub> folds is poorly constrained but appears to be relatively shallow, based on the steep attitude of the L<sub>1</sub> stretching lineation, which is generally oriented perpendicular to the plunge of the associated folds, assuming the folding is cylindrical (Beaumont-Smith and Lavigne, 2008)

The second deformation episode (D<sub>2</sub>) is associated with the development of a prevalent S<sub>2</sub> foliation throughout the map area. D<sub>2</sub> deformation is characterized by shallowly inclined F<sub>2</sub> folds and D<sub>2</sub> thrust faulting. Mesoscopic F<sub>2</sub> folds form prominent structures in the map area: the McLeod Lake synform, the Whitefish bay synform, and another major F<sub>2</sub> synformal structure that can be traced from the Northeast Bay of Herblet Lake trending southwest to Southwest Bay where it wraps around the F<sub>3</sub> Threehouse synform.

During D<sub>3</sub> macroscopic northeast-trending broad, open F<sub>3</sub> folds formed in the Snow Lake area. Locally, an associated axial-planar fabric developed as a spaced cleavage. Decimetre-scale F<sub>3</sub> folds are well developed in Burntwood Group turbidite, whereas the volcanic rocks contain only a local, weak S<sub>3</sub> foliation.

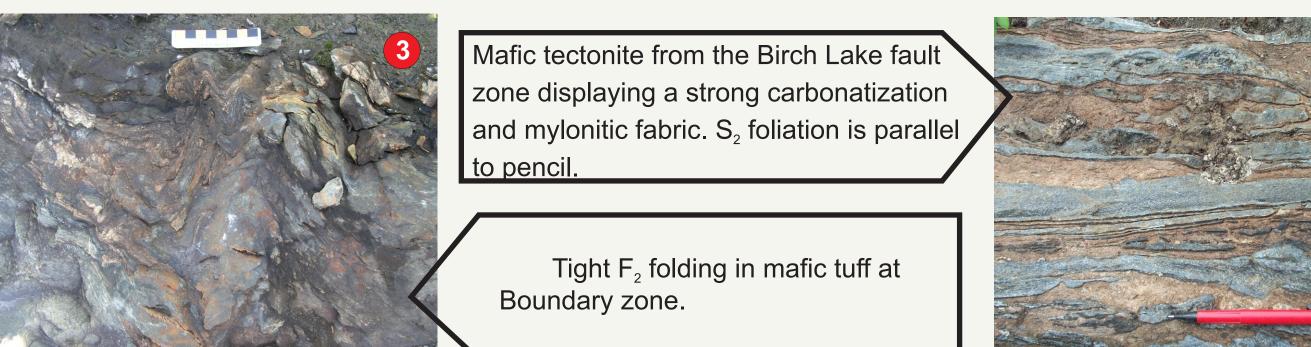
East-trending F<sub>4</sub> folds overprint northeast-trending F<sub>3</sub> structures, resulting in a dome-and-basin configuration just north of the map area. No cleavage associated with the F₄ folding has been identified within the map area

Beaumont-Smith, C.J. and Lavigne, J. 2008: Structural geology and gold metallogenesis of the New Britannia mine area, Snow Lake, Manitoba (NTS 63K16); in Report of Activities 2008, Manitoba Science, Technology, Energy and Mines, Manitoba Geological Survey, p. 7–17. Kraus, J. 1998: Structural and metamorphic studies in the Snow Lake area. Trans-Hudson Orogen, Manitoba, central Canada; Ph.D. thesis, University of New Brunswick, Fredericton, New Brunswick, 229 p. Kraus, J. and Williams, P.F. 1999: Structural development of the Snow Lake allochthon and its role in the evolution of the southeastern Trans-Hudson Orogen in Manitoba. central Canada: Canadian Journal of Earth Sciences, v. 36, no. 11, p. 1881–1899.



Northeast plunging streching lineation 2) developed as quartz rodding in Missi Group meta-arenite on S<sub>2</sub> surface.

Isoclinal F<sub>4</sub> fold in felsic volcaniclastic rocks overprinted by a strong S<sub>2</sub> foliation.



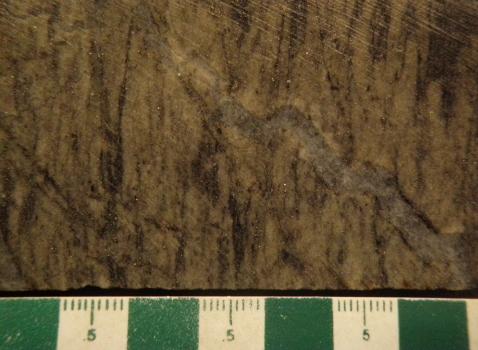
### Gold mineralization



Surface exposure of No. 3 zone



Narrow quartz veins in mineralized (arsenopyrite) and highly silicified gabbro sill at oillowed basalt and mafic volcaniclastics. the Moon Gertie zone.



Thin late quartz vein cross-cutting strongly silicified and foliated rhyodacite rock is mineralized with fine-grained arsenopyrite; Boundary zone.

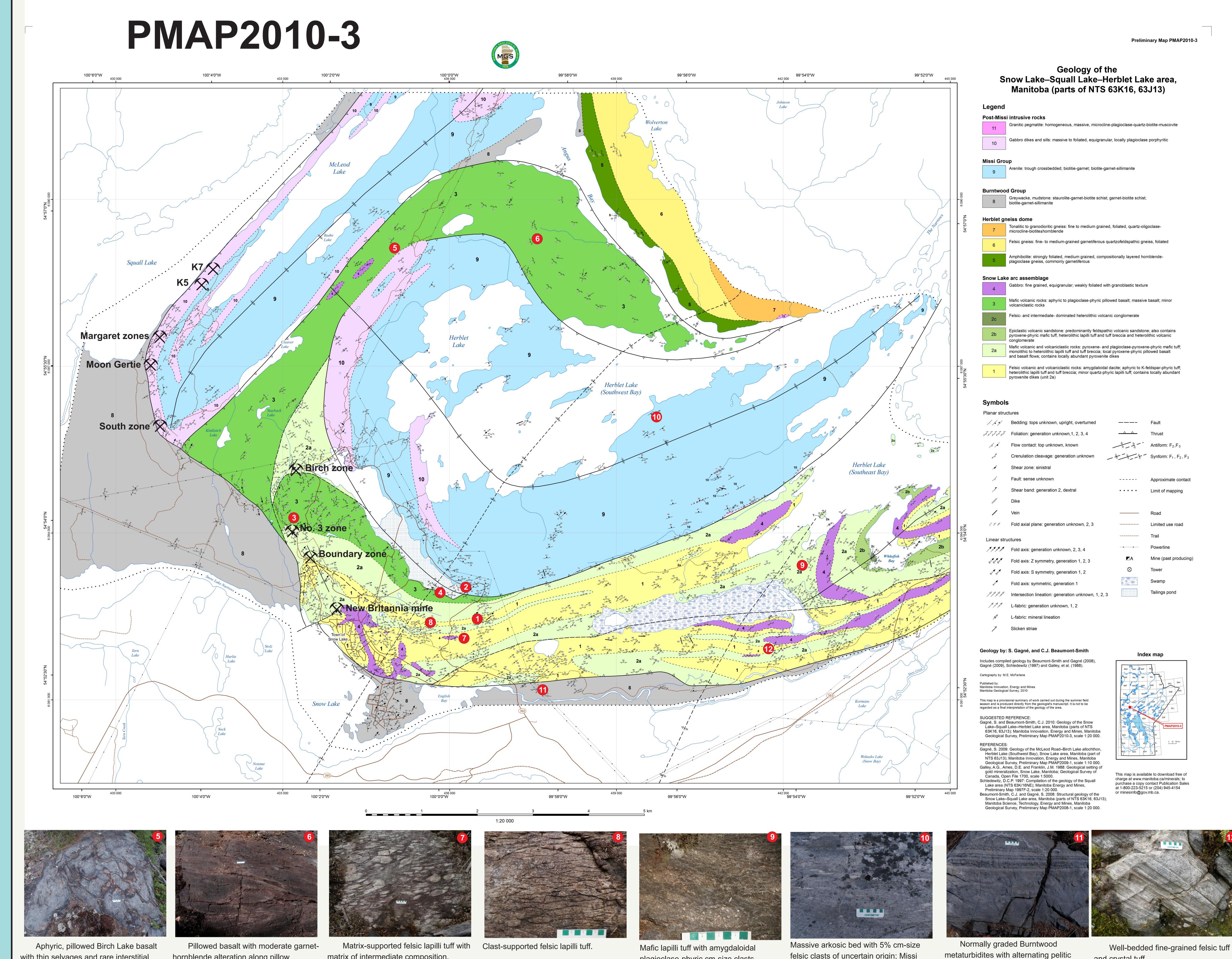


Late barren quartz veins infilling boundinaged and strongly silicified mineralized horizon (medium-grey), within a Bt-Ms-Carb altered host rock; New Britannia mine.

There is a strong spatial association between gold mineralization and the McLeod Road Thrust. Gold mineralization is located in the hangingwall of the thrust and is hosted by a wide variety of rock types. Two common characteristics are the hosting of mineralization within high-strain zones and the location of gold mineralization along or adjacent to lithological

The quartz-vein-hosted gold mineralization and its alteration are overprinted by F<sub>2</sub> folding, which constrains the age of gold emplacement to pre- to syn-D<sub>2</sub>. However, there are also textural relationships that suggest later gold mineralization stages during D<sub>2</sub>. Our current interpretation constrain the initial gold mineralization to syn-D<sub>2</sub> or older accompanied by significant local remobilization during D<sub>2</sub>.

Gold mineralization in the New Britannia mine horizon is hosted by quartz and quartz-carbonate replacements and veins. There appears to be a systematic change in the style of mineralization from east to west. The New Britannia deposit is characterized by massive quartz-carbonate replacement of potassium-altered mafic volcaniclastic rocks adjacent to the contact with felsic volcanic rocks. To the west, the deposits become more closely associated with quartz veins. The Boundary zone contains both quartz replacement and quartz-vein-associated gold mineralization focused at the contact of felsic volcanic and mafic volcaniclastic rocks. Farther west, the No. 3 zone comprise mineralization associated with quartz veins while the Birch zone consists of a fine quartz stockwork.



felsic clasts of uncertain origin; Miss

Group.

plagioclase-phyric cm-size clasts.