

MOLSON LAKE DOMAIN	ISLAND LAKE DOMAIN	BERENS RIVER DOMAIN
<p>Proterozoic</p> <p>Gabbro and diabase dykes, massive (Molson swarm)</p> <p>Granite, leucocratic, biotite-bearing, massive to porphyritic</p> <p>Diabase dykes, massive to foliated</p> <p>Archean</p> <p>Gabbro, minor related granitoid rocks</p> <p>Biotite granulite, tonalite; leucocratic, massive to foliated (2181 ± 89 Ma); Gh - hornblende granulite, tonalite; Gk - megacrystic granulite</p> <p>Quartz diorite, tonalite; leucocratic, strongly foliated; Dh - diorite, quartz diorite, biotite + hornblende-bearing + volcanic and/or sedimentary inclusions</p> <p>Hayes River Group</p> <p>Gabbro, sillstone, intermediate to felsic, biotite-bearing, minor chert</p> <p>Felsic volcanic rocks, massive and fragmental, applied to porphyritic and related intrusive rocks</p> <p>Intermediate to mafic flows and related amphibolite, mafic to felsic, locally gneissic or mylonitic (2265-25 Ma; 3170 ± 68 Ma)</p> <p>Felsic augen gneiss</p> <p>Intermediate layered gneiss with granitoid phases, derived in part from quartzofelsitic sediment; Ni - leucocratic layered gneiss; Nm - mesocratic to melanocratic layered gneiss with granitoid phases, derived in part from mafic volcanic rocks</p> <p>Mafic magmatic gneiss with granitoid phases, in part derived from diorite and gabbro</p>	<p>Granite, leucocratic, biotite-bearing, massive to porphyritic</p> <p>Diabase dykes, massive to foliated</p> <p>Island Lake Group</p> <p>Polysitic conglomerate, minor sandstone</p> <p>Gabbro, massive</p> <p>Biotite granulite, tonalite; leucocratic, massive to foliated (1.2703 Ma - minimum age); Gh - hornblende granulite, tonalite; Gk - megacrystic granulite; Gx - pegmatite</p> <p>Quartz diorite, tonalite; leucocratic, strongly foliated (2683 ± 64 Ma; 2153 ± 48 Ma); Dh - diorite, quartz diorite, biotite + hornblende-bearing + volcanic and/or sedimentary inclusions</p> <p>Felsic augen gneiss</p> <p>Intermediate layered gneiss with granitoid phases, derived in part from quartzofelsitic sediment; Ni - mesocratic to melanocratic layered gneiss with granitoid phases, derived in part from mafic volcanic rocks</p> <p>Mafic magmatic gneiss with granitoid phases, in part derived from diorite and gabbro</p>	<p>Granite, leucocratic, biotite-bearing, massive to porphyritic</p> <p>Diabase dykes, massive to foliated</p> <p>Biotite granulite, tonalite; leucocratic, massive to foliated (2181 ± 89 Ma); Gh - hornblende granulite, tonalite; Gk - megacrystic granulite</p> <p>Quartz diorite, tonalite; leucocratic, strongly foliated (2683 ± 64 Ma; 2153 ± 48 Ma); Dh - diorite, quartz diorite, biotite + hornblende-bearing + volcanic and/or sedimentary inclusions</p> <p>Felsic augen gneiss</p> <p>Intermediate layered gneiss with granitoid phases, derived in part from quartzofelsitic sediment; Ni - mesocratic to melanocratic layered gneiss with granitoid phases, derived in part from mafic volcanic rocks</p> <p>Mafic magmatic gneiss with granitoid phases, in part derived from diorite and gabbro</p>

Reference for U-Pb zircon age:
 1. 89 Emmanovics and Wainess, 1993
 2. 89 Emmanovics and Wainess, 1993

Reference for Pb-Pb zircon age:
 1. 1999 Moore et al., 1999
 2. 1999 Moore et al., 1999
 3. 1999 Moore et al., 1999
 4. 1999 Moore et al., 1999
 5. 1999 Moore et al., 1999
 6. 1999 Moore et al., 1999

Reference for K/Ar mineral ages:
 1. 1978 Loomis, 1981
 2. 1978 Wainess et al., 1978
 3. 1978 Wainess et al., 1978
 4. 1978 Wainess et al., 1978
 5. 1978 Wainess et al., 1978
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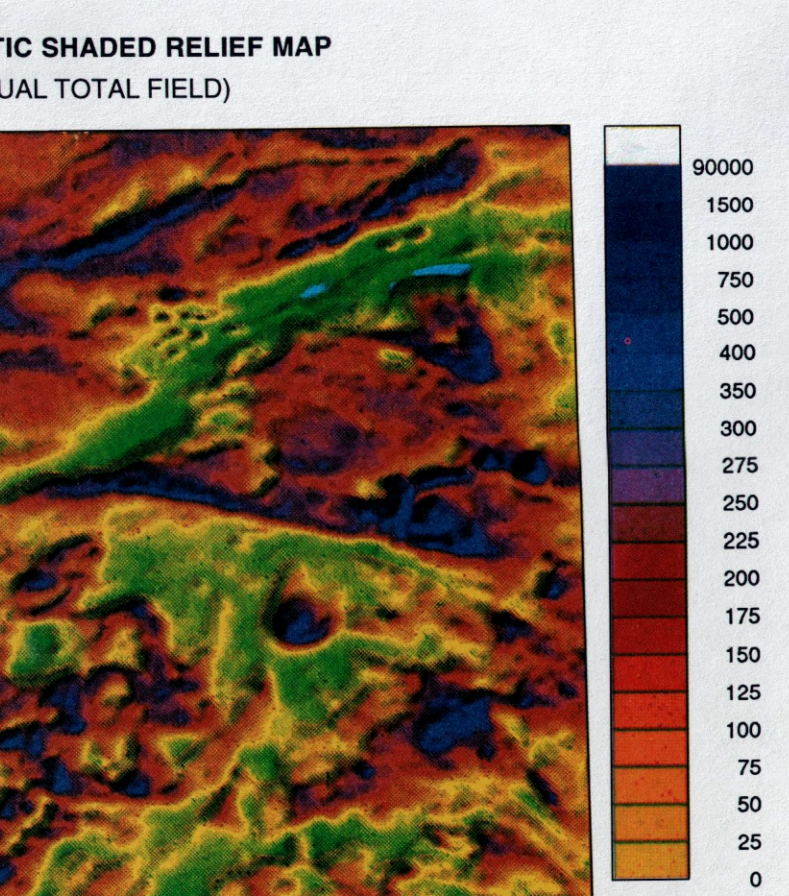
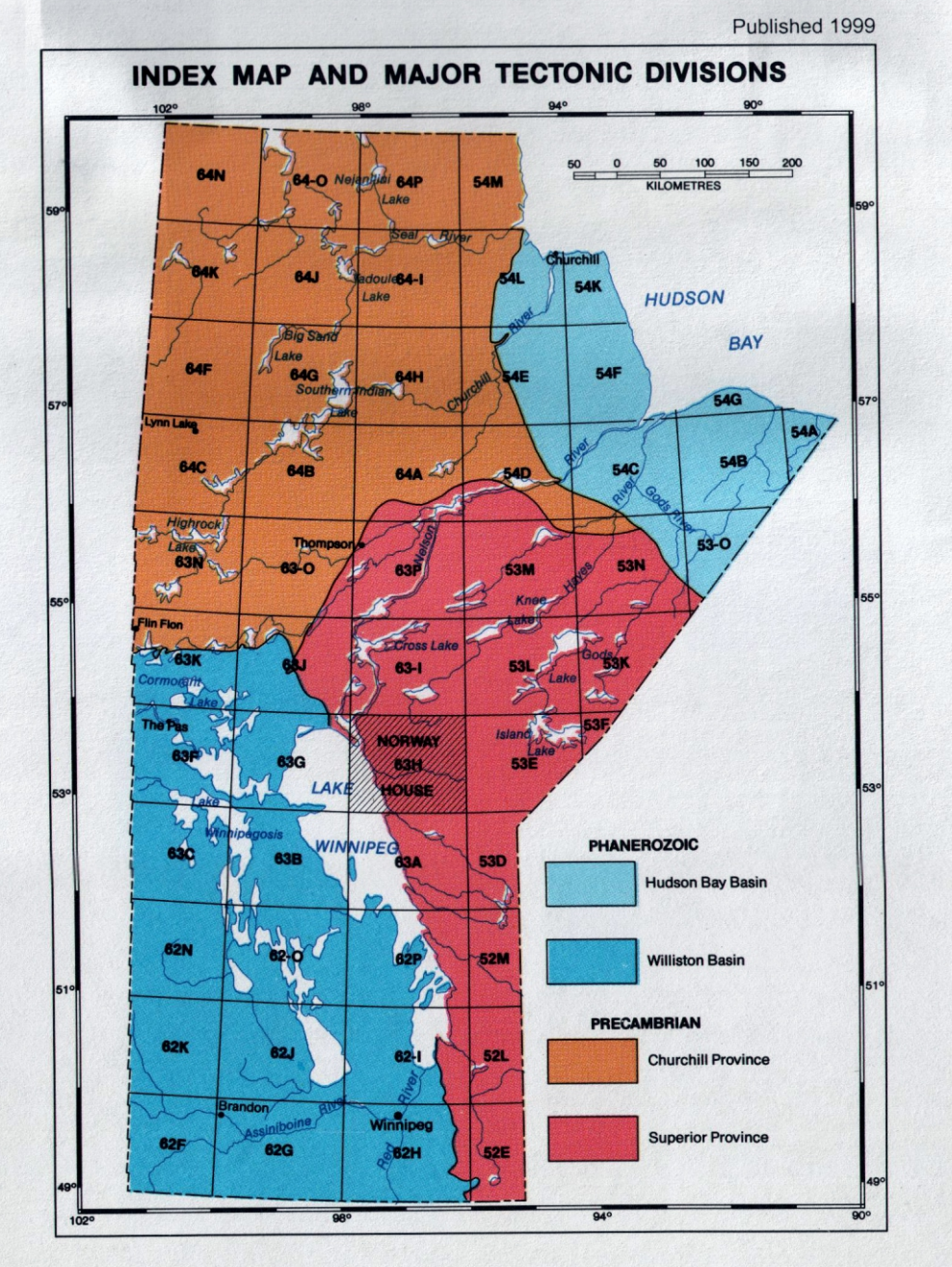
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 5. 1978 Wainess et al., 1978
 6. 1978 Wainess et al., 1978

BEDROCK GEOLOGY COMPILATION MAP SERIES

NORWAY HOUSE
 NTS 63H



MARGINAL NOTES

The Norway House area (NTS 63H) in the Archean Superior Province encompasses parts of the Molson Lake domain to the north (Molson and Island Lake domains) and Berens River subprovince (Berens River domain) to the south (Cory and Wainess, 1993). These domains are terranes that are distinguished by significantly different amounts of supracrustal relative to intrusive rock. Supracrustal rocks constitute 20% of the granitic-greenstone schist subprovince but only 4% of the plutonic Berens River subprovince (Emmanovics and Davison, 1976). Differences in the relative abundance of supracrustal vs. intrusive rocks have been attributed to vertical tectonic movement of domains of different crustal level. The Berens River domain is a major fault (Ayles, 1978). Molson and Berens River domains are relatively more deeply eroded and thus contain less supracrustal (cover) rock compared to Island Lake domain (Emmanovics and Davison, 1976). The fact that rock types, structural patterns and geochronologic data in the various domains are comparable is consistent with their being part of a unified superterrane in the northwestern part of the Superior Province prior to the Mesoproterozoic (Uchi-Sachup-Berens River superterrane; Thurston et al., 1991). However, recent studies show distinctive patterns in Nd-Sm isotopic systematics for several domains in the region (T. Skulski, pers. comm., 1999). These data are interpreted to indicate that each domain evolved along different paths in the Mesoproterozoic prior to amalgamation into a superterrane. Thus the Sm-Nd isotopic variations may indicate that differences between the domains are not simply a reflection of different crustal level, but may be due to pre-supracrustal variations in their lithologic components. Contacts between domains are highly deformed. The Ponask Lake-Stevenson Lake supracrustal belt at the conjunction of the Molson Lake and Island Lake domains is moderately to intensely attenuated and faulted. The contact between the Island Lake and Berens River domains is marked by a mylonite zone (up to 2.4 km wide) that is coincident with a conspicuous aeromagnetic discontinuity (Emmanovics and Davison, 1976).

Sachup subprovince spans approximately 1 billion years of earth history and contains the oldest Superior Province rocks (3.5 Ga Paleoproterozoic), and the products of two subsequent orogenies (3.0-2.9 Ga Wapitigowan and 2.8-2.7 Ga Laurentian; Emmanovics and Wainess, 1993). Supracrustal rocks in Sachup subprovince consist of coarse-grained, arc-related, igneous sequences and related sedimentary rocks dated at 2.7-3.0 Ga (Corty et al., 1992; Thurston et al., 1991). The oldest supracrustal rocks that are basal members in many supracrustal sequences in Ontario (Corty and Wood, 1996). Island Lake domain supracrustal rocks 2683 ± 64 Ma volcanic-derived amphibole at Stevenson Lake (2; Emmanovics and Wainess, 1993), which is the oldest dated rock in the Norway House area. 3047 ± 2 Ma detrital zircon in sandstone at Cross Lake (Corty et al., 1992) attest to Paleoproterozoic magmatism that preceded existing supracrustal belts.

The Ponask Lake-Stevenson Lake greenstone belt extends for approximately 65 km along the northwest margin of Island Lake domain, and extends a further 150 km to the east, through the Island Lake area (NTS 53E) into Ontario. The supracrustal rocks in the Stevenson Lake area are significantly more attenuated compared to those at Island Lake. However, lithologic units and the stratigraphic sequence are consistent with those established at Island Lake, a volcanosedimentary sequence (Hayes River Group) is unconformably overlain by fluvial/alluvial conglomerates and sandstone turbidites (Island Lake Group; Weber, 1982; Neale, 1984; Gilbert, 1985). The basal Hayes River Group unit at Stevenson Lake (w) consists of gabbro, followed by massive basalt and related gabbro up to 0.5 km thick, equivalent rocks at Island Lake have an estimated thickness of 0.8 to 1.2 km. Ultramafic lenses and small plugs occur sporadically within the volcanic rocks. Massive to fragmental felsic volcanic and related intrusive rocks (w) up to 120 m thick and a 180 m wide pyroxene-sillstone unit (w) apparently overlie the mafic volcanic section. The sedimentary rocks display normal size grading, and contain sporadic garnet and cordierite. Polymorphic conglomerates (w) extend for at least 45 km through the centre of the Stevenson Lake supracrustal belt, the 300 m wide unit is interpreted to occupy the core of a major syncline (Gilbert, 1985). The conglomerate, which consists mainly of felsic plutonic fragments with subordinate volcanic, sedimentary and quartz clasts, is lithologically similar to the basal Island Lake Group conglomerate at Island Lake (Weber, 1981). The unit rests unconformably on post-Hayes River Group tonalite and quartz diorite (D) in the east part of Stevenson Lake and northwest part of Island Lake (NTS 53E).

Layered gneisses (N) are interpenetrated with plutonic rocks throughout the Molson and Island Lake domains, and are a subordinate component of the many granitoid Berens River domain. The magmatic gneisses are characterized by alternating leucocratic and amphibolite layers and minor sills, and are interpreted as derived from (a) supracrustal rocks that were intruded by pervasive granitoid phases, and (b) tonalite that was intruded by diabase and gabbro; precursors probably include both Mesoproterozoic and Neoproterozoic igneous gneisses. Synmetamorphic and late K-metasomatism associated with granitoid/igneous intrusions are characteristic of the gneisses. The layered gneisses are classified as leucocratic (N), melanocratic (Nm) mesocratic (Nn) and mafic magmatic, derived in part from diorite and gabbro (N). Contacts between layered gneiss and gneissite belts are typically tectonic, whereas contacts within gneisses are gradational (Thurston et al., 1991).

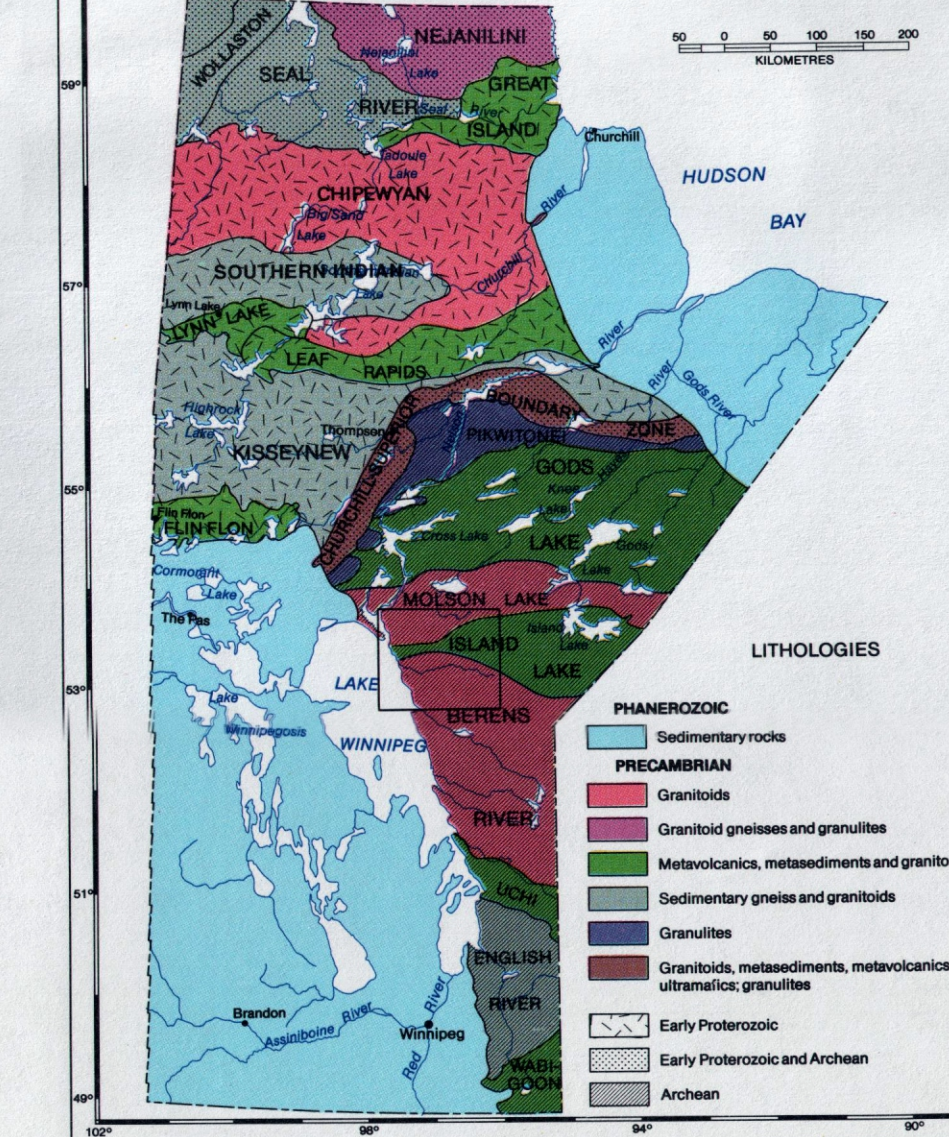
The four main granitoid rock suites in the Norway House area, possibly correlative with suites 1 to 4 in the Sachup and Berens River subprovinces in Ontario (Thurston et al., 1991), are in order of decreasing age:

- quartz diorite to granitoid felsic augen gneiss (Ga), interpreted as coeval with tonalitic phases in layered gneiss (N);
 - foliated quartz diorite and tonalite (D), diorite (Dh);
 - massive to foliated granulite (G, Gh, Gk, Gx);
 - massive granite (G).
- In Ontario, suite 1 includes 2.95 Ga orthogneiss (Corty and Ayles, 1984) that is approximately coeval with the oldest supracrustal rocks in the Sachup and Berens River subprovinces; suite 2 (2.87-2.71 Ga) and suite 3 (2.70 Ga; Thurston et al., 1991) postdate volcanic rocks in the Ponask-Stevenson Lake belt, whereas suite 4 (ca. 2.65 Ga; Mezger et al., 1990) postdates both the volcanic rocks and overlying Island Lake Group fluvial sedimentary rocks.
- Quartz diorite to granitoid felsic augen gneiss (Ga, suite 1) is characterized by mafic magmatic segregations, quartz aggregates and stringers, and mafic xenoliths and schlieren assumed to be derived from supracrustal rocks (Emmanovics, 1973). Quartz diorite and tonalite (D, suite 2), which also contain supracrustal enclaves and related schlieren, are moderately to strongly foliated, and mylonitic toward contacts with gneissite belts; the granitoid rocks contain abundant epidote in the area east of Lake Winnipeg. Mafic intrusions of suite 2 (foliated diorite, Dh) are abundant in the Island Lake domain and contiguous parts of Molson Lake domain; the mafic intrusions are assumed coeval with quartz diorite (D). Predominantly massive granulite of suite 3 underlies over 60% of the Norway House area (NTS 63H); subdivisions of this suite are based on mafic mineral type (biotite, G, hornblende, Gh) and texture (with microlite megacrysts, Gx; pegmatite, Gk). The granulite rocks are relatively free of alteration, in contrast to more recrystallized quartz diorite/tonalite (D). Massive, coarse-grained to pegmatitic granite dykes and small plutons (G, suite 4) occur within and at the margins of older granitoid terranes. This suite includes muscovitic, leucocratic and 2-mica granite; in Ontario, related pegmatite dykes, emplaced in major shear zones, are commonly fractionated and contain rare metal minerals (Thurston et al., 1991). The late (suite 4) granites have been interpreted as anatectic products of metasedimentary origin, and are commonly located at major shear zones and/or domain boundaries (Thurston et al., 1991; Breaks and Osman, 1989).
- Three mafic intrusive suites postdate Hayes River Group supracrustal rocks. Sporadic mafic dykes and small plugs (gabbro (B), altered to amphibolite) are largely confined to the granulite terrane (G) of Molson Lake domain, and are interpreted as pre-Island Lake Group. Massive to foliated diabase dykes, which occur mostly in Sachup River subprovince, postdate granitoid rocks except late granite rocks (4), and are equalled with post-Island Lake Group metabasites in the Island Lake area (Weber et al., 1982). These dykes, which are also altered to amphibolite, are generally conformable with host rock foliation trends. North to northeast trending, massive gabbro and diabase dykes are common in the marginal, northwest part of Superior Province (Escobedo and Maclell, 1978); within the Norway House map sheet, all but one of these dykes occur in Sachup subprovince. The Proterozoic Molson dykes (1884 ± 2 Ma; Heaman et al., 1986) display fresh igneous textures, and locally contain olivine, pyroxene and/or plagioclase phenocrysts.

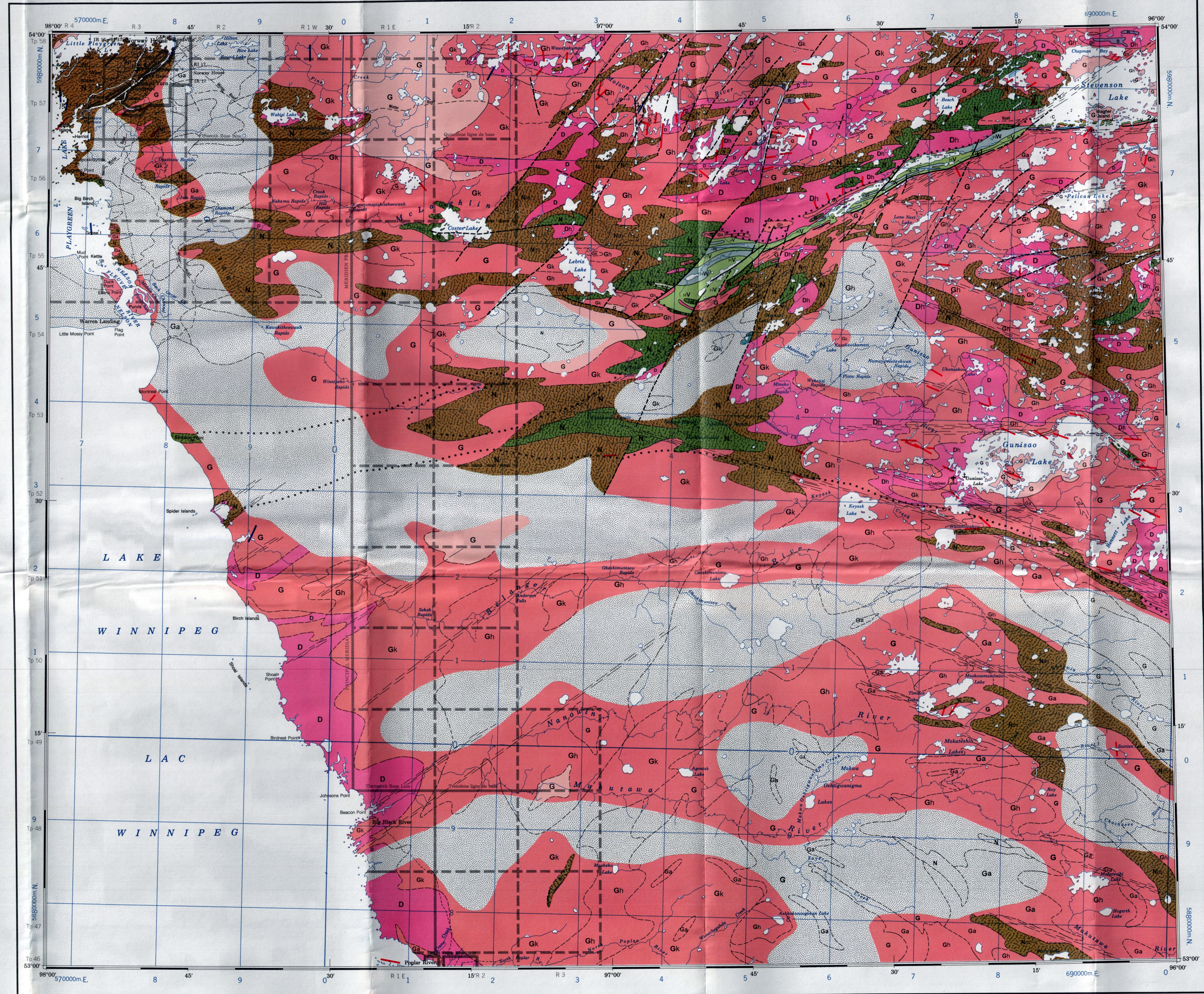
Multiple metamorphic events have affected supracrustal and early plutonic rocks in the Norway House area. U/Pb geochronology studies of metamorphic garnets have documented three Neoproterozoic high grade metamorphic events in the western Sachup subprovince between 2.74 and 2.64 Ga (Mezger, 1990). Subsequent regional greenschist facies retrogression is locally intensified in shear zones, and in attenuated supracrustal rocks in the Ponask-Stevenson Lake belt. The influence of the Hudsonian orogeny at the margin of the northwestern Superior Province is documented by K/Ar and Rb/Sr age determinations that record the resetting of Archean rock (tonalite) ages in a zone that extends for up to 350 km southeast of the Churchill-Superior boundary (Emmanovics and Wainess, 1993). The effects of this ca. 1.80-1.70 Ga Paleoproterozoic metamorphism were probably very similar to, and indistinguishable from, the preceding Neoproterozoic retrogressive metamorphism.

Major, belt-parallel folds in the Ponask Lake-Stevenson Lake belt are accompanied by a regional foliation, which was in turn deformed in a later fold event (Gilbert, 1985). Warping of major fold axial planes may have been due to the emplacement of granitoid plutons, which were associated in Ontario with localized homing of regionally foliated rocks (Thurston et al., 1991). Northeast and west-northeast trending shear zones that transect the map area are interpreted as coeval with vertical displacement at domain-boundary faults, which is assumed to postdate the folding.

PRINCIPAL GEOLOGICAL DOMAINS



Every possible effort has been made to ensure that the information presented on this map is accurate. However, the Province of Manitoba Energy and Mines do not assume liability for any errors that may occur. References are included for users wishing to verify critical information.

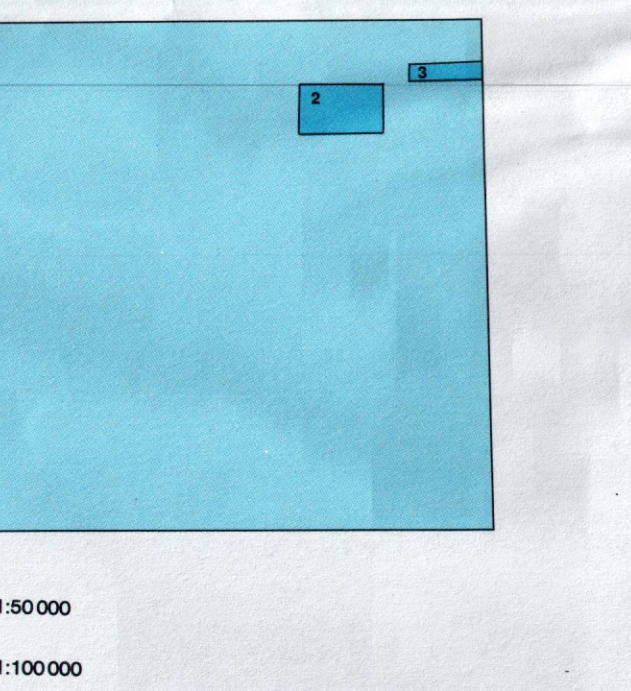


Scale 1:250 000
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 Printed 1999

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1. Emmanovics, I.F., 1971. Norway House and Grand Rapids map area. Geological Survey of Canada, Map 5-1972 (accompanies Paper 72-29); scale 1:250 000.
 2. Weber, W., 1986. Ponask Lake-Manitoba Energy and Mines, Preliminary Map 1986-1-1; scale 1:20 000.
 3. Gilbert, H.P., 1985. Wagner Island; Manitoba Energy and Mines, Preliminary Map 1985-1-5; scale 1:20 000.

Approximate mean declination (1999) for centre of map (Downloaded by G. Lindal)

SYMBOLS

- Geological boundary (approximate, interpreted)
- Fault (defined, assumed)
- Domain boundary
- Area of title or no outcrop
- Sample locality for U-Pb or Pb-Pb zircon age determination
- Shear zone

AEROMAGNETIC SHADED RELIEF MAP (RESIDUAL TOTAL FIELD)

